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22

23 **Key Points:**

- 24 ● Biological processes and fluvial inputs drive seasonal variations in surface dissolved trace metals
25 (dTMs) on Celtic Sea.
- 26 ● Dissolved Al, Cd, Ni, Cu, and Zn concentrations and metal:P ratios at depth are controlled by water

27 mass mixing.

28 ● Mediterranean Outflow Waters provide a strong imprint on dTM concentrations and metal:P ratios in
29 the Northeast Atlantic Ocean.

30

31 **Abstract**

32 We report the seasonal distributions of dissolved zinc (dZn), nickel (dNi), copper (dCu), cadmium
33 (dCd), aluminum (dAl), and nutrients on the Northeast Atlantic continental margin (Celtic Sea). Surface
34 dissolved trace metal (dTM) and nutrient variations were mainly regulated by the seasonal cycling of
35 biological processes. The dTM (especially for dCu and dZn) and nutrient stoichiometry on the continental
36 shelf was additionally affected by fluvial input. Nutrients and dTMs at depth on the continental slope were
37 determined by water mass mixing driven by ocean circulation, without invoking local remineralization
38 process. The Mediterranean Outflow Waters are especially important for delivering Mediterranean-sourced
39 dTMs into the Northeast Atlantic Ocean and drive dTM:nutrient kinks at a depth of ~ 1000 m. These results
40 highlight the importance of riverine input, seasonality of primary production and ocean circulation on the
41 distributions of nutrients and nutrient-like dTMs in temperate continental margins, which could further
42 affect local biological carbon pump.

43

44 **Plain Language Summary**

45 Nutrients and nutrient-like dissolved trace metals (dTMs; dZn, dNi, dCd, dCu), as well as their
46 correlations are important for understanding biogeochemical cycles in paleo- and modern oceans. Here we
47 show that dTM and nutrients on the continental shelf of Northeast Atlantic continental margin was
48 determined by the balance between river-derived materials and local biogeochemical processes. Seasonal
49 variations in surface dTMs and nutrients on the continental slope were mainly regulated by biological
50 processes. On the other hand, dTMs and nutrients at depth were controlled by the mixing of water masses
51 with different dTM:nutrient stoichiometry. The most important changes in dTM:nutrient ratios (kinks) at a
52 depth of ~ 1000 m were closely associated with the Mediterranean Outflow Waters. These findings suggest
53 that seasonal variations in dTMs and nutrients on the Northeast Atlantic continental margin were

Commented [AB1]: Is there much evidence for Al, Cd, Ni, Cu, and Zn restricting productivity near continental margins?

Commented [CX2R1]: Should remove these words?

54 dynamically controlled by external sources (terrestrial materials from rivers), local biogeochemical
55 processes, and ocean circulation. Short-term and long-term temporal variations of these factors may
56 consequently change the availability of nutrients and micronutrients to continental margins, which are
57 important transition zones for the marine carbon cycle.

58

59 1. Introduction

60 Dissolved (< 0.2 µm) bioactive trace metals (dTM) including zinc (Zn), nickel (Ni), copper (Cu) and
61 perhaps cadmium (Cd), are important micronutrients in marine systems. Zinc, Ni, and Cu are involved in
62 enzymatic processes required from phytoplankton growth (La Fontaine et al., 2002; Mahaffey et al., 2014;
63 Twining & Baines, 2013) and Cd is taken up as a divalent metal (Horner et al., 2013). Low supply of dTMs,
64 therefore, could potentially affect marine ecosystem structure and functioning (Lohan & Tagliabue, 2018;
65 Morel & Price, 2003). Due to their biological uptake, the vertical distributions of dissolved Zn (dZn), Ni
66 (dNi), Cu (dCu), and Cd (dCd) resemble those of nutrients (Nitrate+Nitrite (TN), phosphate (P), and silicic
67 acid (Si)) (Bruland et al., 2014). These dTMs typically exhibit seasonal depleted concentrations in surface
68 waters due to phytoplankton uptake (Moore et al., 2013; Morel & Price, 2003) and elevated levels at depth
69 due to remineralization of sinking organic particles (Boyd et al., 2017). Nutrient and bioactive dTMs,
70 therefore, show significant positive correlations for example the dCd-P relationships in global oceans
71 (Boyle et al., 1976; Middag et al., 2018; Wu & Roshan, 2015; Xie et al., 2015), dZn-P, dNi-P, and dCu-Si
72 correlations in the South Atlantic Ocean (Wyatt et al., 2014) and the Southern Ocean (e.g., Janssen et al.,
73 2020; Saito et al., 2010).

74 The linear relationships between bioactive dTMs and nutrients usually show pronounced changes in
75 slopes (kinks), e.g., at a P concentration of ~ 1.3 µM for the Cd-P correlation (de Baar et al., 1994; Cullen,
76 2006). The origin of such kinks (especially the Cd:P kink) has been debated over the last decades. Some
77 hypotheses point towards deeper regeneration of Cd relative to P (Boyle, 1988; Roshan & DeVries, 2021),
78 or enhanced Cd uptake due to the limitation of bio-essential elements in surface waters (Cullen, 2006;
79 Sunda & Huntsman, 2000). Kinks are also hypothesized to be associated with the chemical replacement
80 between Co, Zn, and Cd in carbonic anhydrase in phytoplankton (Morel et al., 1994; Price & Morel, 1990)
81 or changing bioavailability of Cd through organic complexation (Bruland, 1992). Recent studies
82 demonstrated that the mixing of water masses with different Cd:P ratios could be pivotal for the observed
83 kinks (Baars et al., 2014; Middag et al., 2018; Xie et al., 2015). In addition, external sources of trace metals
84 such as continental inputs and dust deposition can also affect bioactive dTM distributions in the ocean, such
85 inputs can be traced by dissolved aluminum (dAl) (Han et al., 2008; Measures & Edmond, 1988; Menzel

Commented [GM3]: Yes I think you could remove this.

Commented [ML4R3]: I agree it can be removed but as talking about biological process I would just have one sentence stating- Zn, Ni and Cu are involved in enzymatic process required from phytoplankton growth (add a few refs) and Cd is taken up as divalent metal (ref).

Commented [CX5R3]: done

Commented [AB6]: Excess Cu also toxic

Commented [CX7R6]: Yes.

86 Barraqueta et al., 2018). The ongoing debate over the drivers underpinning bioactive dTM distributions
87 requires further investigation into the co-occurrences of nutrients and dTMs, better informing this debate is
88 crucial for our understanding biogeochemical cycles in paleo – and modern oceans.

89 Continental margins with their shelves and slopes are junctions between terrestrial systems and open
90 oceans. The disproportionately high primary production and particulate organic carbon export make
91 continental margins important transition zones for the marine carbon cycle (Muller-Karger et al., 2005;
92 Simpson & Sharples, 2012). Here, we report on the seasonal distributions of dTMs and nutrients on the
93 Northeast (NE) Atlantic continental margin (Celtic Sea), which is characterized by large seasonal variations
94 of biological activities (Birchill et al., 2017), complex bathymetry and dynamic water circulation (Fig. 1a).
95 High sampling and seasonal resolution (Fig. 1b) offer a unique opportunity to determine the influence of
96 terrestrial inputs, biogeochemical processes, and ocean circulation on the seasonal variations of bioactive
97 dTMs and their relationships with nutrients.

98

99 **2. Methods**

100 Samples were collected on board the *RRS Discovery* during three different seasons: an autumn cruise
101 in November 2014 (DY018), a spring cruise in April 2015 (DY029), and a summer cruise in July 2015
102 (DY033). We conducted one transect on the continental shelf of Celtic Sea, from station Site A near the
103 Bristol channel to station CS2 near the shelf break (Fig. 1b) during each season. Two off-shelf transects
104 were conducted along a canyon (stations Fe01 - Fe07, Fe15) and a spur (stations Fe08 - Fe14) during each
105 season.

106 Seawater samples for dTM analyses were collected following GEOTRACES protocols (Cutter et al.,
107 2017). Samples were filtered immediately upon collection using a 0.2 µm filter capsule (Acropack). Trace
108 metals were pre-concentrated at GEOMAR using an automated system (SC-4 DX SeaFAST pico; ESI) and
109 analyzed by high-resolution inductively coupled plasma mass spectrometry (HR-ICP-MS, Thermo Fisher
110 Element XR) as outlined in Rapp et al. (2017). Dissolved Al concentrations were analyzed by a
111 spectrofluorometer (Cary Eclipse) as per Hydes & Liss (1976). ~~Short lived radium (Ra) isotopic activities,
112 ²²³Ra (half life, $t_{1/2}$ = 11.4 days) and ²²⁴Ra (half life, $t_{1/2}$ = 3.66 days), were counted from large volume~~

113 ~~samples using a Radium Delayed Coincidence Counter (Moore, 2008). Radium isotope activities here are~~
114 ~~reported in excess of activity supported by their parent isotopes in the water column.~~ Nutrients (P, Si, and
115 TN) were measured on board using techniques described in Woodward & Rees (2001), according to the
116 International GO-SHIP nutrient manual recommendations (Hydes et al., 2010). Further details of sampling
117 and analyses can be found in the [supporting information](#) and Rusiecka et al. (2018).

118

119 3. Results and discussion

120 Dissolved Cd, Zn, Ni, and Cu exhibited nutrient-like vertical distributions during all seasons on the
121 NE Atlantic continental margin (Fig. 1c, Fig. 2, Fig. 3), with the lowest concentrations observed in surface
122 waters due to biological utilization, and elevated concentrations at depth ascribed to remineralization
123 (Bruland et al., 2014; Moore et al., 2013). The canyon and spur transects showed comparable vertical
124 profiles for dTMs and nutrients in each season and were combined as a slope transect. In contrast to
125 relatively constant of bioactive dTM and nutrient concentrations at depth on the continental slope, surface
126 bioactive dTM and nutrient levels on Celtic Sea showed pronounced seasonal variations (Table S2). Kinks
127 in dTM:nutrient (e.g., dZn:P and dCu:P) ratios were observed on the continental slope at a depth of ~ 100
128 m and ~1000 m (Fig. S2). The shelf transect (Fig. 3) showed higher dCu, dNi, dZn, and Si concentrations
129 than the slope transect (Fig. 1c, Fig. 2), accompanied by varying dTM: nutrient correlations.

130

131 3.1 Biological influence on seasonal variations in surface dTM along the continental slope

132 A seasonal mixed layer (SML) with a depth of ~ < 100 m from spring to autumn occurs on the
133 continental slope (Supporting information: Hydrography). A reduction in dTM and nutrient levels in the
134 SML (especially at depths < 30 m) was observed between April 2015 and July 2015, due to biological
135 utilization and water column stratification (Birchill et al., 2017) (Table S2). The total drawdown of surface
136 (depth < 30 m) dTMs and nutrients on the slope from April to July were: ~~dAl 2.68 ± 2.25 nM~~, dCd 79.5 ±
137 45.1 pM, dCu 0.08 ± 0.11 nM, dNi, 0.67 ± 0.31 nM, dZn 0.10 ± 0.25 nM, P 0.35 ± 0.11 μM, TN 6.71 ±
138 1.83 μM, and Si 2.44 ± 0.47 μM (Table S2). The removal of dCd, dNi, P, and TN over this period was higher,
139 and drawdown of Cu was lower than reported for this region between January 1994 and June 1995 (TN 3.9

Commented [AB8]: Is this true? Identical suggests that at each corresponding sampling depth the concentrations determined remained within the analytical uncertainty of the measurements throughout April, July and November. At least in surface waters there must be some seasonal variation?

Maybe rephrase this, something like "showed comparable distributions of dTMs and nutrients in each season."

Commented [CX9R8]: done

Commented [AB10]: I would remove of example and state the ones you are referring to precisely,

Commented [CX11R10]: There are too many to present: dCu and dZn with all nutrients, dCd with TN and Si, dNi and Si.

Commented [AB12]: For Zn also in deeper waters around 2500m?

In the SI might be good include profiles of the dTM:P ratio so the reader can easily see this

Commented [CX13R12]: There are not so many samples at depth > 2500 m. The slopes of linear regression models are the ratios (also summarised at Table S3).

Commented [ML14]: I can't follow this when I look at table S2 (need to add units to the table). How calculated- is this the mean and then std deviation? Do you integrate values within each depth range? As the standard variation is very large- e.g. for Zn and Ni higher than the drawdown- what is the main point here- Not really clear the drawdown- know it is occurring but perhaps you have too wide a region.

Commented [CX15R14]: Yes, these are mean ± sd values. The standard variations are very large, similar to previous studies. The ratios are calculated using mean values.

140 $\pm 3.9 \mu\text{M}$, P $0.46 \pm 0.51 \mu\text{M}$, Si $2.2 \pm 3.7 \mu\text{M}$, Cd $39 \pm 61 \text{ pM}$, Cu $0.34 \pm 0.58 \text{ nM}$, Ni $0.5 \pm 1.1 \text{ nM}$) (Cotté-
 141 Krief et al., 2002) and March – June 1987 (Cd $30 \pm 12 \text{ pM}$) (Kremling & Pohl, 1989). The decrease in dTM
 142 and nutrient concentrations was the consequence of phytoplankton uptake in summer, and the overall
 143 “uptake” ratio of phytoplankton normalized to P was:

$$(N_{19}Si_7P_1)_{1000}Al_{7.5}Ni_{1.9}Zn_{0.34}Cu_{0.23}Cd_{0.23}$$

144 The surface dTM and nutrient concentrations on the slope increased from July to November assuming
 145 no significant annual changes in seawater chemistry at the same month, accompanied by comparable
 146 correlations between dTMs/nutrients and AOU at AOU > 0 (Fig. S3). Therefore, the surface dTM and
 147 nutrient variations from summer to autumn is additionally affected by the remineralization of sinking
 148 organic particles. On the other hand, the increase in surface dTM and nutrient concentrations from
 149 November to April was attributed to resupply from subsurface waters (Birchill et al., 2017). Using
 150 concentration differences between July and November (Table S2), we estimated the apparent
 151 “remineralization” ratio of dTMs and nutrients normalized to P as:

$$(N_{15}Si_6P_1)_{1000}Al_{4.9}Ni_{1.9}Zn_{0.56}Cu_{0.49}Cd_{0.28}$$

154 and the “winter mixing” ratio estimated from the concentration differences between November and April
 155 observations as:

$$(N_{21}Si_{7.5}P_1)_{1000}Al_{5.4}Ni_{1.8}Zn_{0.19}Cu_{0.06}Cd_{0.19}$$

157 The estimated Zn:P and Cu:P ratios between “uptake” and “remineralization” varied by as much as a
 158 factor 2, but the “winter mixing” ratios were as much as an order of magnitude lower, probably due to a
 159 relatively limited seasonal variations and large concentration ranges observed of both metals (Table S2).
 160 The N:P, Si:P, Ni:P, and Cd:P ratios were relatively constant, indicating a close association of Ni and Cd
 161 with biological processes in surface waters across all seasons. The observed “uptake”, “remineralization”,
 162 and “winter mixing” ratios are close to the overall dTM:P ratios in the SML (depth of $< \sim 100 \text{ m}$) ($(P_{1000}Ni_{1.53}$
 163 $\pm 0.08Zn_{0.37} \pm 0.07Cu_{0.20} \pm 0.03Cd_{0.21} \pm 0.01)$) (Table S3). The dTM:P ratio here is broadly consistent with the extended
 164 Redfield ratio of phytoplankton cultures ($(N_{16}P_1)_{1000}Zn_{0.80}Cu_{0.38}Cd_{0.21}$) (Ho et al., 2003). Therefore, the
 165 positive correlations between dTMs and nutrients in the SML on the continental slope across all seasons
 166 (Fig. S4) generally reflected the seasonal cycling of biological uptake in summer, remineralization of

Commented [GM16]: This red bit and then next one are referencing literature values. I guess they are not strictly necessary, but also it a shame to remove them. Reviewers generally like to see references to literature values and if you don't have them you get complaints...

Commented [ML17R16]: I agree but these are expressed as % but the data from this paper as concentrations...

Commented [DR18R16]: I agree with Maeve, this is expressed in % whilst our values in actual concentrations

Commented [AB19]: Aluminium, abiotic removal processes too?

Commented [CX20R19]: Should dAl be removed in this section to avoid controversy?

Commented [ML21]: This is great and perhaps all you need- just state drawdown and then this ratio if need to keep paper short

Commented [ML22]: You could get pulled up here as we don't have the Nov after July i.e. Nov 2014 and July 2015- need to put a caveat in here. I think it is reasonable to assume would be similar the next Nov but need to state this

Commented [DR23R22]: Yes, I agree, this needs to be clarified that it's an assumption of the same cycle

Commented [CX24R22]: done

Commented [ML25]: Be careful here as can't use AOU in surface waters as have negative values- remove this from fig. S7 and be specific where you have evidence of remineralisation.

Commented [CX26R25]: Done. I removed AOU < 0 in the figure

Commented [ML27]: Agree with this

Commented [AB28]: Extra aluminium maybe from external dust supply. I think there was a Saharan dust episode that ...

Commented [CX29R28]: dAl removed from this section?

Commented [CX30]: Antony: Maybe you need to consider propagating the uncertainty for these ratios? ...

Commented [CX31R30]: done

167 organic particles in autumn, followed by winter mixing.

168 3.2 Additional fluvial inputs of dTMs on the shelf

169 Seasonal cycling of biological processes also affected the dTM and nutrient distributions on the shelf
170 of the NE Atlantic Ocean. Using station CCS (central Celtic Sea) as an example (Fig. S4), surface dTM and
171 nutrient concentrations decreased from April to July due to phytoplankton uptake. Subsurface (depths > 50
172 m) nutrient and dCd levels increased from April to November due to remineralization of sinking organic
173 particles (Birchill et al., 2017; Lohan & Tagliabue, 2018). In contrast, subsurface dCu and dNi gradually
174 decreased from spring to autumn, possibly reflecting the impact of external sources and/or water mass
175 mixing. ~~Unlike similar dTM:P ratios at all stations on the slope,~~ The overall dTM:P ratios on the shelf varied
176 greatly between sampling locations (Fig. S5). These variations were accompanied by increasing dTM
177 concentrations with decreasing distance offshore, suggesting the dTM stocks on the shelf were additionally
178 supplied by external sources. Benthic sediments were likely not an important source for the enhanced dTMs,
179 since dTM concentrations did not change significantly with $^{223}\text{Ra}_{\text{xs}}$ and $^{224}\text{Ra}_{\text{xs}}$ activities (Fig. S6).

180 Instead, an increasing salinity with distance offshore (Fig. 3) and strong negative correlations between
181 dTMs and salinity suggest dTM (especially for dCu and dZn) concentrations (Fig. S7) were augmented by
182 a dTM-rich low-salinity endmember, e.g., riverine input from the British Isles through Irish Sea and/or the
183 Bristol channel (Achterberg et al., 1999; Kremling & Hydes, 1988). Based on the correlations between
184 subsurface (depth of 50 – 200 m to exclude surface biological activities) dTMs, nutrients, and salinity in
185 April when significant correlations were observed, the low-salinity endmember at salinity of 33.5 (the
186 typical salinity of the Irish Sea and the Bristol Channel) were calculated as: dCd, 144 ± 59 pM; dCu, 13.0
187 ± 1.2 nM; dNi, 4.45 ± 1.11 nM; dZn, 22.2 ± 3.2 nM; TN, 9.45 ± 3.85 μM ; P, 0.80 ± 0.23 μM ; and Si, 12.9
188 ± 1.8 μM . These values are comparable with the observed dTM concentrations in the low-salinity waters
189 around the British Isles (Kremling & Hydes, 1988). The enrichment of dCu and dZn relative to P in the
190 low-salinity endmember produced gradually decreasing dCu:P and dZn:P ratios with increasing distance
191 stretching from Site A to CS2 (Fig. S9). Fluvial input was not a major source of dCd and dNi, thereby
192 resulting in increasing dCd:P and dNi:P ratios with offshore distance. ~~At station CS2, the dTM:P, TN:P, and~~
193 ~~Si:P ratios were close to those on the continental slope.~~

Commented [DR32]: Out of curiosity, do you have a suggestion why the ratios of DTMs changed whilst macronutrients remained the same?

Commented [CX33R32]: Nutrients were also changed

Commented [AB34]: I agree with the broader point that any seasonal cycling on the shelf is superimposed on top of a salinity and hence TM concentration gradient. But looking at Fig S8 the uptake and remineralisation picture is less clear than described here?

Yes, there is clear surface drawdown for dCd, looks like dFe and macro nutrients, dZn, surface draw down yes but remineralisation less clear. dCu concentrations are relatively homogenous throughout the water column for each month.

For both dNi and dCu concentrations throughout the water column decrease each month. Does the salinity change and maybe indicate a water mass signal? As the Cu and Ni

Commented [CX35R34]: Yes. dCd is controlled by local biological processes. dCu and dZn are determined by

Commented [AB36]: Salinity increases with distance offshore?

Commented [CX37R36]: Yes. It's a mistake. Salinity increases with offshore distance.

Commented [AB38]: In Fig S12 the correlations for Cu and Zn are convincing but not so for Cd and Ni. Is it possible that

Commented [CX39R38]: Yes. I think Cu and Zn are mainly controlled by external sources, while Cd was mainly

Commented [AB40]: I think for this part you really do need to consider the estuarine behaviour, this assumes a linear

Commented [ML41]: Could this be why you see larger variation in the offshore transects? E.g. large ranges of

Commented [CX42R41]: This section refers to the shelf region, while the former section refers to the slope region

Commented [DR43]: So on-shelf waters (with fluvial input) were lower in dCd and dNi than the slope surface waters

Commented [CX44R43]: No. Just because P is somewhat contributed by fluvial input while Cd and Ni rarely. So the

194 ~~The subsurface salinity at station Site A gradually increased from ~34.9 in April to ~35.3 in November~~
195 ~~(Fig. 3), suggesting decreasing fluxes of riverine waters or an increasing intrusion of North Atlantic waters.~~
196 ~~For instance, the water flow of River Severn, the longest river of the British Isles, shows decreasing flows~~
197 ~~from winter to summer and autumn [website, Open WIMS data].~~ The low-salinity endmember shows the
198 highest dCu and dZn concentrations in July and lowest values in November (Fig. S7), possibly reflecting
199 seasonal variations in the endmember dTM concentrations. This phenomenon can be additionally explained
200 by the enhanced influence of remineralization at stations away from the fluvial source, evidenced by the
201 elevated subsurface dTM and nutrient levels at station CCS relative to other stations in autumn (Fig. S7).
202 Therefore, the distributions of dTMs and nutrients as well as their correlations on the continental shelf were
203 balanced by riverine input and the seasonal cycling of biogeochemical processes.

204

205 3.3 Water mass mixing drive metal:P kinks at depth

206 The waters on the NE Atlantic continental slope between the SML and potential density of 27.62 kg
207 m⁻³ (depth ~ 1000 m) are characterized by the presence of East North Atlantic Central Waters (ENACW),
208 Mediterranean Outflow Waters (MOW), and Sub-Arctic Intermediate Waters (SAIW) (Rusiecka et al., 2018)
209 (Fig. S8). The increasing MOW contribution with depth is accompanied by increasing dTM and nutrient
210 concentrations. ~~At depths of 950–1050 m with the highest MOW contribution (~60%), waters showed~~
211 ~~strongly elevated dAl (20.1 ± 1.5 nM) (Table S4).~~ Water columns with potential density > 27.62 kg m⁻³ are
212 characterized by a gradually decreasing MOW contribution, and increasing contributions of Labrador Sea
213 Water (LSW) and North East Atlantic Deep Waters (NEADW) (Fig. S8). The dTM and nutrient
214 concentrations continuously increased with depth, showing dCd of ~ 350 pM, dCu of ~ 2.2 nM, dNi of ~ 5
215 nM, dZn of ~ 2.7 nM in bottom waters (Table S2). [These concentrations are similar to the reported deep
216 dCd (310 ± 26 pM), dNi (4.1 ± 0.4 nM), and dCu (1.56 ± 0.33 nM) values for this region (Cotté-Krief et
217 al., 2002) and consistent with reported deep water dTM and nutrient levels in the North Atlantic Ocean
218 (Achterberg et al., 2021; Saager et al., 1997).]

219 No apparent kinks were identified for dCd:P (260 ± 3 μmol mol⁻¹), dNi:P (1.94 ± 0.04 mmol mol⁻¹),
220 and dZn:P (2.26 ± 0.04 mmol mol⁻¹) in waters > 100 m (Table S3). These ratios here are similar to those

Commented [AB45]: I think more likely this will be driven by regional ocean-shelf circulation, would be worth reading anything by Sharples or Jo Hopkins relating to the circulation of the Celtic Sea

Commented [CX46R45]: Possibly this sentence could be removed.

Commented [ML47]: I would leave this in

Commented [CX48]: Antony: You could summarise this, my suggestion would be to hold off on the deleting until you have addressed the comments and then we may be able to edit our way under the word count. It is nice to place the values in context.

221 reported for the North Atlantic Ocean with dCd:P of $278 \pm 3 \mu\text{mol mol}^{-1}$, dNi:P of $2.01 \pm 0.08 \text{ mmol mol}^{-1}$,
222 and dZn:P of $1.77 \pm 0.16 \text{ mmol mol}^{-1}$ at P of 0.5 – 1.5 μM (GEOTRACES Intermediate Data Product Group,
223 2021; Middag et al., 2018; Roshan & Wu, 2015). The lack of dCd:P kink also agrees with the linear dCd –
224 P relationship at $P < 1.3 \mu\text{mol kg}^{-1}$ (de Baar et al., 1994; Cullen, 2006; Frew & Hunter, 1992; Middag et al.,
225 2018). In contrast, the dCu:P ratio on the slope of NE Atlantic continental margin increased from 0.31 mmol
226 mol^{-1} at 100 – 1000 m to 2.78 mmol mol^{-1} at depths > 1000 m and the dAl concentrations showed
227 pronounced variations with increasing P levels (Fig. S2). Considering the small variations in surface dCu
228 concentrations (Table S2) and that Al is not a bio-essential element, changes in subsurface dCu:P and dAl:P
229 ratios with water depth should reflect physical (e.g., water mass mixing) rather than biological processes.

230 We estimated the elemental composition of each water mass using a three-step calculation (Table S4).
231 (1) At potential density < 27.62 kg m^{-3} (depths < ~ 1000 m), endmember MOW concentrations were
232 calculated from the positive linear correlations with dTM concentrations (Fig. S9). (2) Nutrients and dTMs
233 contributed by MOW were removed. At potential density > 27.62 kg m^{-3} , the residual dTMs were
234 contributed by LSW, NEADW, and SAIW, where the endmember SAIW concentrations were calculated
235 from the significant negative linear relationship between corrected SAIW contribution and corrected dTM
236 concentrations (Fig. S10). The endmember NEADW concentrations were evaluated at LSW < 1% (Fig.
237 S11a). (3) Finally, the endmember concentrations of LSW and ENACW were estimated by removing the
238 contributions of MOW, SAIW, and NEADW (Fig. S11b and c).

239 All dTMs and nutrients showed significant correlations with percentage contributions of LSW and
240 ENACW at the final step, despite uncertainties propagating during each step in the calculations. The
241 predicted dTM concentrations, reconstructed by direct multiplication of water mass fractions with their
242 endmember values, illustrate almost identical values with the observed concentrations with very low
243 residuals (Fig. S12). The calculated endmember concentrations of NEADW agree with deep water (> 4000
244 m) concentrations in the NE Atlantic Ocean (GEOTRACES Intermediate Data Product Group, 2021; Liu
245 & Tanhua, 2021), where NEADW is a persistent feature (van Aken, 2000a; García-Ibáñez et al., 2015, 2018;
246 Reinthaler et al., 2013). Therefore, our estimations on the apparent endmember concentrations of water
247 masses are robust to show their relative chemical compositions.

Commented [GM49]: Is there a reason you switch from "." to "-"?

Commented [CX50R49]: Here dCd-P means the correlation between dCd and P, not their ratio.

Commented [AB51]: Is this shown? What type of variation?

Commented [CX52R51]: Twisted variations, Fig. S2

Commented [DR53]: Please do not put Ni and Al in one sentence here. They are driven by different processes. Whilst DCu was probably actively taken up whilst DAl passively scavenged. DAl also had maxima whilst DCu did not

Commented [CX54R53]: We are talking about subsurface samples here. They are combined because they both reflect water mass mixing.

Commented [AB55]: Not sure despite is the correct word. You could make a statement about the uncertainty and state a value(s)?

Commented [CX56R55]: It is difficult to state the uncertainty values

Commented [GM57]: If you remove the red bit above, then it may be tricky to make this conclusion?

Commented [CX58R57]: Antony: Could you add any observed end member values to Table S4 from the GEOTRACES IDP?

Commented [CX59R57]: done

248 The correlations between reconstructed dTMs and nutrients corresponded to the observed results, that
249 no kinks were observed for the correlations between dCd, dNi, dZn, and P, while dCu:P and dAl:P ratios
250 changed sharply at a density of 27.62 kg m^{-3} (depth $\sim 1000 \text{ m}$) (Fig. 4). The dTM:AOU and nutrient:AOU
251 ratios changed abruptly at depths of $\sim 1000 \text{ m}$ and $\sim 2000 \text{ m}$ across all seasons (Fig. S3), coinciding with
252 the variations of water mass fractions from MOW+SAIW+ENACW at $100 - 1000 \text{ m}$ to
253 MOW+NEADW+LSW at $> \sim 1000 \text{ m}$. Therefore, the AOU variations at depth mostly reflect physical
254 processes (e.g., water mass mixing) rather than local biological processes. All these observations indicate
255 that subsurface dTMs and nutrients and their ratios on the NE Atlantic continental margin are mainly
256 controlled by water mass mixing driven by ocean circulation with local remineralization making a minor
257 contribution.

258 3.4 The impact of MOW on the dTM distributions on the NE Atlantic Ocean

259 The dTM:nutrient and dTM:AOU kinks at $\sim 1000 \text{ m}$ and $\sim 2000 \text{ m}$ are closely related to the maximum
260 and diminished occurrence of MOW (Fig. S2, Fig. S3), probably ascribed to the distinctive dTM and
261 nutrient stoichiometry of MOW relative to other water masses (Table S4). For instance, MOW shows much
262 higher dAl:P but lower dCu:P ratios than LSW and NEADW, thus creating kinks of dCu:P and dAl:P ratios
263 at the maximum occurrence of MOW. Therefore, MOW provides an important imprint on the dTM
264 distributions on the NE Atlantic continental slope.

265 The MOW is formed in the Mediterranean Sea and spreads across the NE Atlantic Ocean at $\sim 500 -$
266 1500 m towards the Bay of Biscay and further along the shelf break of Celtic Sea (van Aken, 2000b; Price
267 et al., 1993). ~~The occurrence of MOW in the NE Atlantic Ocean can be observed in elevated dAl~~
268 ~~concentrations and salinity (Measures et al., 2015; Middag et al., 2022; Rolison et al., 2015) at depths of~~
269 ~~900–1400 m.~~ The significant correlations between dTMs and salinity (Fig. S13) demonstrate that dTMs
270 in the MOW core were predominantly controlled by the conservative isopycnal mixing between MOW and
271 lower salinity water masses (e.g., SAIW with similar density range to MOW; Johnson & Gruber, 2007)
272 during ocean circulation, rather than removal by scavenging. Specifically, the dAl, dZn, and dNi
273 concentrations of the MOW core decreased with decreasing salinity, suggesting the saline MOW is a net
274 source to deliver Mediterranean-sourced Al, Zn, and Ni into the NE Atlantic Ocean (Middag et al., 2022).

Commented [DR60]: If MOW is Zn and Ni source then why don't we see maxima in DZn and Dni profiles similarly to DAl?

275 This finding is similar to the long-distance transport of anthropogenic Pb from MOW to the NE Atlantic
276 continental margin (Rusiecka et al., 2018).

277 Furthermore, the minor seasonal variations of dTMs and nutrients in the MOW core possibly reflect
278 seasonal cycles of water mass circulation along the continental slope. Due to the wind-driven processes
279 (Roque et al., 2019), the influence of SAIW in the NE Atlantic Ocean declines in autumn. Hence, higher
280 MOW signals (e.g., higher salinity) were observed along the continental slope in November with respect to
281 April and July. Accordingly, the MOW core on the slope showed slightly higher dAl and dZn but lower
282 dCd concentrations in November than those in April and July (Fig. S13).

283

284 4. Conclusions

285 Our findings illustrate the seasonal variations of surface dTMs and nutrients were associated with
286 biological processes on the continental margin of the NE Atlantic Ocean. Surface dTM concentrations on
287 the shelf were also influenced by a low-salinity endmember, likely fluvial materials from the British Isles.
288 Therefore, temperate shelf sea ecosystems can be influenced by local biological processes and external
289 sources, where riverine inputs play an essential role to deliver terrestrial dTMs to the ocean. The dTM
290 concentrations and metal:P ratios at depth in the slope region can be explained by water mass mixing driven
291 by ocean circulation without invoking local remineralization. Specifically, the long-distance transportation
292 of MOW delivers Mediterranean-sourced dTMs (e.g., dAl, dZn, and Ni) into the NE Atlantic Ocean and
293 drives dAl:P and dCu:P kinks at a potential density of $\sim 27.62 \text{ kg m}^{-3}$ (depth $\sim 1000 \text{ m}$) along the NE Atlantic
294 continental slope. Future climate change driven changes in dust inputs into the Mediterranean and water
295 mass characteristics in the subpolar gyre, therefore, will have consequences for nutrient stoichiometry and
296 the biological carbon cycles in the NE Atlantic Ocean.

297

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300 expeditions and Malcolm Woodward and Carolyn Harrys for the macronutrient data. This project was
301 funded by the UK Natural Environment Research Council (NE/K001973/1 (E. A. and M. G.),

Commented [CX61R60]: Because water mass will evolve during transportation

Commented [DR62]: Do you mean seasonal variations in water mass circulation?

Commented [CX63R62]: Yes

Commented [DR64]: How do you know its not ENACW?

Commented [CX65R64]: See Johnson and Gruber, 2007

Commented [AB66]: this is hard to see in the figure S23

Commented [CX67R66]: slightly higher

302 NE/K001779/1 (M. L.), NE/K002023/1 (A. A.), and NE/L501840/1 (A. B.)).

303 Open Research

304 Data are held at the British Oceanographic Data Centre (<http://www.bodc.ac.uk/>).

305

306 References

- 307 Achterberg, E. P., Colombo, C., & van den Berg, C. M. G. (1999). The distribution of dissolved Cu, Zn, Ni,
308 Co and Cr in English coastal surface waters. *Continental Shelf Research*, 19(4), 537–558.
309 [https://doi.org/10.1016/S0278-4343\(98\)00093-4](https://doi.org/10.1016/S0278-4343(98)00093-4)
- 310 Achterberg, E. P., Steigenberger, S., Klar, J. K., Browning, T. J., Marsay, C. M., Painter, S. C., et al. (2021).
311 Trace Element Biogeochemistry in the High-Latitude North Atlantic Ocean: Seasonal Variations and
312 Volcanic Inputs. *Global Biogeochemical Cycles*, 35(3), e2020GB006674.
313 <https://doi.org/10.1029/2020GB006674>
- 314 van Aken, H. M. (2000a). The hydrography of the mid-latitude northeast Atlantic Ocean: I: The deep water
315 masses. *Deep Sea Research Part I: Oceanographic Research Papers*, 47(5), 757–788.
316 [https://doi.org/10.1016/S0967-0637\(99\)00092-8](https://doi.org/10.1016/S0967-0637(99)00092-8)
- 317 van Aken, H. M. (2000b). The hydrography of the mid-latitude Northeast Atlantic Ocean: II: The
318 intermediate water masses. *Deep Sea Research Part I: Oceanographic Research Papers*, 47(5), 789–
319 824. [https://doi.org/10.1016/S0967-0637\(99\)00112-0](https://doi.org/10.1016/S0967-0637(99)00112-0)
- 320 de Baar, H. J. W., Saager, P. M., Nolting, R. F., & van der Meer, J. (1994). Cadmium versus phosphate in
321 the world ocean. *Marine Chemistry*, 46(3), 261–281. [https://doi.org/10.1016/0304-4203\(94\)90082-5](https://doi.org/10.1016/0304-4203(94)90082-5)
- 322 Baars, O., Abouchami, W., Galer, S. J. G., Boye, M., & Croot, P. L. (2014). Dissolved cadmium in the
323 Southern Ocean: Distribution, speciation, and relation to phosphate. *Limnology and Oceanography*,
324 59(2), 385–399. <https://doi.org/10.4319/lo.2014.59.2.0385>
- 325 Birchill, A. J., Milne, A., Woodward, E. M. S., Harris, C., Annett, A., Rusiecka, D., et al. (2017). Seasonal
326 iron depletion in temperate shelf seas. *Geophysical Research Letters*, 44(17), 8987–8996.
327 <https://doi.org/10.1002/2017GL073881>
- 328 Boyd, P. W., Ellwood, M. J., Tagliabue, A., & Twining, B. S. (2017). Biotic and abiotic retention, recycling
329 and remineralization of metals in the ocean. *Nature Geoscience*, 10(3), 167–173.
330 <https://doi.org/10.1038/ngeo2876>
- 331 Boyle, E. A., Sclater, F., & Edmond, J. M. (1976). On the marine geochemistry of cadmium. *Nature*,
332 263(5572), 42–44. <https://doi.org/10.1038/263042a0>
- 333 Boyle, Edward A. (1988). Cadmium: Chemical tracer of deepwater paleoceanography. *Paleoceanography*,
334 3(4), 471–489.
- 335 Bruland, K. W. (1992). Complexation of cadmium by natural organic ligands in the central North Pacific.
336 *Limnology and Oceanography*, 37(5), 1008–1017. <https://doi.org/10.4319/lo.1992.37.5.1008>
- 337 Bruland, K. W., Middag, R., & Lohan, M. C. (2014). Controls of Trace Metals in Seawater. In *Treatise on*
338 *Geochemistry* (pp. 19–51). Elsevier. <https://doi.org/10.1016/B978-0-08-095975-7.00602-1>
- 339 Cotté-Krief, M.-H., Thomas, A. J., & Martin, J.-M. (2002). Trace metal (Cd, Cu, Ni and Pb) cycling in the

340 upper water column near the shelf edge of the European continental margin (Celtic Sea). *Marine*
341 *Chemistry*, 79(1), 1–26. [https://doi.org/10.1016/S0304-4203\(02\)00013-0](https://doi.org/10.1016/S0304-4203(02)00013-0)

342 Cullen, J. T. (2006). On the nonlinear relationship between dissolved cadmium and phosphate in the modern
343 global ocean: Could chronic iron limitation of phytoplankton growth cause the kink? *Limnology and*
344 *Oceanography*, 51(3), 1369–1380. <https://doi.org/10.4319/lo.2006.51.3.1369>

345 Cutter, G., Casciotti, K., Croot, P., Geibert, W., Heimbürger, L.-E., Lohan, M., et al. (2017). Sampling and
346 Sample-handling Protocols for GEOTRACES Cruises. Version 3, August 2017. *GEOTRACES*
347 *Standards and Intercalibration Committee*.

348 Frew, R. D., & Hunter, K. A. (1992). Influence of Southern Ocean waters on the cadmium–phosphate
349 properties of the global ocean. *Nature*, 360(6400), 144–146. <https://doi.org/10.1038/360144a0>

350 García-Ibáñez, M. I., Pardo, P. C., Carracedo, L. I., Mercier, H., Lherminier, P., Ríos, A. F., & Pérez, F. F.
351 (2015). Structure, transports and transformations of the water masses in the Atlantic Subpolar Gyre.
352 *Progress in Oceanography*, 135, 18–36. <https://doi.org/10.1016/j.pocean.2015.03.009>

353 García-Ibáñez, M. I., Pérez, F. F., Lherminier, P., Zunino, P., Mercier, H., & Tréguer, P. (2018). Water mass
354 distributions and transports for the 2014 GEOVIDE cruise in the North Atlantic. *Biogeosciences*, 15(7),
355 2075–2090. <https://doi.org/10.5194/bg-15-2075-2018>

356 GEOTRACES Intermediate Data Product Group. (2021). The GEOTRACES Intermediate Data Product
357 2021 (IDP2021). *NERC EDS British Oceanographic Data Centre NOC*.
358 <https://doi.org/10.5285/cf2d9ba9-d51d-3b7c-e053-8486abc0f5fd>

359 Han, Q., Moore, J. K., Zender, C., Measures, C., & Hydes, D. (2008). Constraining oceanic dust deposition
360 using surface ocean dissolved Al. *Global Biogeochemical Cycles*, 22(2).

361 Ho, T.-Y., Quigg, A., Finkel, Z. V., Milligan, A. J., Wyman, K., Falkowski, P. G., & Morel, F. M. M. (2003).
362 The Elemental Composition of Some Marine Phytoplankton1. *Journal of Phycology*, 39(6), 1145–
363 1159. <https://doi.org/10.1111/j.0022-3646.2003.03-090.x>

364 Horner, T. J., Lee, R. B. Y., Henderson, G. M., & Rickaby, R. E. M. (2013). Nonspecific uptake and
365 homeostasis drive the oceanic cadmium cycle. *Proceedings of the National Academy of Sciences*,
366 110(7), 2500–2505. <https://doi.org/10.1073/pnas.1213857110>

367 Hydes, D. J., & Liss, P. S. (1976). Fluorimetric method for the determination of low concentrations of
368 dissolved aluminium in natural waters. *Analyst*, 101(922), 922–931.
369 <https://doi.org/10.1039/an9760100922>

370 Hydes, D. J., Aoyama, M., Aminot, A., Bakker, K., Becker, S., Coverly, S., & Daniel, A. (2010).
371 Determination of dissolved nutrients (N, P, Si) in seawater with high precision and inter-comparability
372 using gas-segmented continuous flow analysers. In *The GO-SHIP Repeat Hydrography Manual: A*
373 *Collection of Expert Reports and Guidelines., IOCCP No 1* (OCPO Publication Series No. 134, version
374 1), 1–87.

375 Janssen, D. J., Sieber, M., Ellwood, M. J., Conway, T. M., Barrett, P. M., Chen, X., et al. (2020). Trace
376 metal and nutrient dynamics across broad biogeochemical gradients in the Indian and Pacific sectors
377 of the Southern Ocean. *Marine Chemistry*, 221, 103773.
378 <https://doi.org/10.1016/j.marchem.2020.103773>

379 Johnson, G. C., & Gruber, N. (2007). Decadal water mass variations along 20°W in the Northeastern
380 Atlantic Ocean. *Progress in Oceanography*, 73(3), 277–295.

381 <https://doi.org/10.1016/j.pocean.2006.03.022>

382 Kremling, K., & Hydes, D. (1988). Summer distribution of dissolved Al, Cd, Co, Cu, Mn and Ni in surface
383 waters around the British Isles. *Continental Shelf Research*, 8(1), 89–105.
384 [https://doi.org/10.1016/0278-4343\(88\)90026-X](https://doi.org/10.1016/0278-4343(88)90026-X)

385 Kremling, K., & Pohl, C. (1989). Studies on the spatial and seasonal variability of dissolved cadmium,
386 copper and nickel in northeast atlantic surface waters. *Marine Chemistry*, 27(1), 43–60.
387 [https://doi.org/10.1016/0304-4203\(89\)90027-3](https://doi.org/10.1016/0304-4203(89)90027-3)

388 La Fontaine, S., Quinn, J. M., Nakamoto, S. S., Page, M. D., Göhre, V., Moseley, J. L., et al. (2002). Copper-
389 dependent iron assimilation pathway in the model photosynthetic eukaryote *Chlamydomonas*
390 *reinhardtii*. *Eukaryotic Cell*, 1(5), 736–757.

391 Liu, M., & Tanhua, T. (2021). Water masses in the Atlantic Ocean: characteristics and distributions. *Ocean*
392 *Science*, 17(2), 463–486. <https://doi.org/10.5194/os-17-463-2021>

393 Lohan, M. C., & Tagliabue, A. (2018). Oceanic Micronutrients: Trace Metals that are Essential for Marine
394 Life. *Elements*, 14(6), 385–390. <https://doi.org/10.2138/gselements.14.6.385>

395 Mahaffey, C., Reynolds, S., Davis, C. E., & Lohan, M. C. (2014). Alkaline phosphatase activity in the
396 subtropical ocean: insights from nutrient, dust and trace metal addition experiments. *Frontiers in*
397 *Marine Science*, 1, 73.

398 Measures, C., & Edmond, J. M. (1988). Aluminium as a tracer of the deep outflow from the Mediterranean.
399 *Journal of Geophysical Research: Oceans*, 93(C1), 591–595.
400 <https://doi.org/10.1029/JC093iC01p00591>

401 Measures, C., Hatta, M., Fitzsimmons, J., & Morton, P. (2015). Dissolved Al in the zonal N Atlantic section
402 of the US GEOTRACES 2010/2011 cruises and the importance of hydrothermal inputs. *Deep Sea*
403 *Research Part II: Topical Studies in Oceanography*, 116, 176–186.
404 <https://doi.org/10.1016/j.dsr2.2014.07.006>

405 Menzel Barraqueta, J.-L., Schlosser, C., Planquette, H., Gourain, A., Cheize, M., Boutorh, J., et al. (2018).
406 Aluminium in the North Atlantic Ocean and the Labrador Sea (GEOTRACES GA01 section): roles of
407 continental inputs and biogenic particle removal. *Biogeosciences*, 15(16), 5271–5286.
408 <https://doi.org/10.5194/bg-15-5271-2018>

409 Middag, R., van Heuven, S. M. A. C., Bruland, K. W., & de Baar, H. J. W. (2018). The relationship between
410 cadmium and phosphate in the Atlantic Ocean unravelled. *Earth and Planetary Science Letters*, 492,
411 79–88. <https://doi.org/10.1016/j.epsl.2018.03.046>

412 Middag, R., Rolison, J. M., George, E., Gerringa, L. J. A., Rijkenberg, M. J. A., & Stirling, C. H. (2022).
413 Basin scale distributions of dissolved manganese, nickel, zinc and cadmium in the Mediterranean Sea.
414 *Marine Chemistry*, 238, 104063. <https://doi.org/10.1016/j.marchem.2021.104063>

415 Moore, C. M., Mills, M. M., Arrigo, K. R., Berman-Frank, I., Bopp, L., Boyd, P. W., et al. (2013). Processes
416 and patterns of oceanic nutrient limitation. *Nature Geoscience*, 6(9), 701–710.
417 <https://doi.org/10.1038/ngeo1765>

418 Moore, W. S. (2008). Fifteen years experience in measuring ²²⁴Ra and ²²³Ra by delayed-coincidence
419 counting. *Marine Chemistry*, 109(3–4), 188–197.

420 Morel, F. M. M., Reinfelder, J. R., Roberts, S. B., Chamberlain, C. P., Lee, J. G., & Yee, D. (1994). Zinc
421 and carbon co-limitation of marine phytoplankton. *Nature*, 369(6483), 740–742.

422 <https://doi.org/10.1038/369740a0>

423 Morel, Francois M. M., & Price, N. M. (2003). The biogeochemical cycles of trace metals in the oceans.
424 *Science*, 300(5621), 944–947.

425 Muller-Karger, F. E., Varela, R., Thunell, R., Luerssen, R., Hu, C., & Walsh, J. J. (2005). The importance
426 of continental margins in the global carbon cycle. *Geophysical Research Letters*, 32(1), L01602.

427 Price, J. F., Baringer, M. O., Lueck, R. G., Johnson, G. C., Ambar, I., Parrilla, G., et al. (1993).
428 Mediterranean Outflow Mixing and Dynamics. *Science*, 259(5099), 1277–1282.
429 <https://doi.org/10.1126/science.259.5099.1277>

430 Price, N. M., & Morel, F. M. M. (1990). Cadmium and cobalt substitution for zinc in a marine diatom.
431 *Nature*, 344(6267), 658–660.

432 Rapp, I., Schlosser, C., Rusiecka, D., Gledhill, M., & Achterberg, E. P. (2017). Automated preconcentration
433 of Fe, Zn, Cu, Ni, Cd, Pb, Co, and Mn in seawater with analysis using high-resolution sector field
434 inductively-coupled plasma mass spectrometry. *Analytica Chimica Acta*, 976, 1–13.
435 <https://doi.org/10.1016/j.aca.2017.05.008>

436 Reinthaler, T., Álvarez Salgado, X. A., Álvarez, M., van Aken, H. M., & Herndl, G. J. (2013). Impact of
437 water mass mixing on the biogeochemistry and microbiology of the Northeast Atlantic Deep Water.
438 *Global Biogeochemical Cycles*, 27(4), 1151–1162. <https://doi.org/10.1002/2013GB004634>

439 Rolison, J. M., Middag, R., Stirling, C. H., Rijkenberg, M. J. A., & de Baar, H. J. W. (2015). Zonal
440 distribution of dissolved aluminium in the Mediterranean Sea. *Marine Chemistry*, 177, 87–100.
441 <https://doi.org/10.1016/j.marchem.2015.05.001>

442 Roque, D., Parras-Berrocá, I., Bruno, M., Sánchez-Leal, R., & Hernández-Molina, F. J. (2019). Seasonal
443 variability of intermediate water masses in the Gulf of Cádiz: implications of the Antarctic and
444 subarctic seesaw. *Ocean Science*, 15(5), 1381–1397.

445 Roshan, S., & DeVries, T. (2021). Global Contrasts Between Oceanic Cycling of Cadmium and Phosphate.
446 *Global Biogeochemical Cycles*, 35(6), e2021GB006952. <https://doi.org/10.1029/2021GB006952>

447 Roshan, S., & Wu, J. (2015). Cadmium regeneration within the North Atlantic. *Global Biogeochemical*
448 *Cycles*, 29(12), 2082–2094. <https://doi.org/10.1002/2015GB005215>

449 Rusiecka, D., Gledhill, M., Milne, A., Achterberg, E. P., Annett, A. L., Atkinson, S., et al. (2018).
450 Anthropogenic Signatures of Lead in the Northeast Atlantic. *Geophysical Research Letters*, 45(6),
451 2734–2743. <https://doi.org/10.1002/2017GL076825>

452 Saager, P. M., de Baar, H. J. W., de Jong, J. T. M., Nolting, R. F., & Schijf, J. (1997). Hydrography and local
453 sources of dissolved trace metals Mn, Ni, Cu, and Cd in the northeast Atlantic Ocean. *Marine*
454 *Chemistry*, 57(3), 195–216. [https://doi.org/10.1016/S0304-4203\(97\)00038-8](https://doi.org/10.1016/S0304-4203(97)00038-8)

455 Saito, M. A., Goepfert, T. J., Noble, A. E., Bertrand, E. M., Sedwick, P. N., & DiTullio, G. R. (2010). A
456 seasonal study of dissolved cobalt in the Ross Sea, Antarctica: micronutrient behavior, absence of
457 scavenging, and relationships with Zn, Cd, and P. *Biogeochemistry*, 7(12), 4059–4082.
458 <https://doi.org/10.5194/bg-7-4059-2010>

459 Simpson, J. H., & Sharples, J. (2012). *Introduction to the physical and biological oceanography of shelf*
460 *seas*. Cambridge University Press.

461 Sunda, W. G., & Huntsman, S. A. (2000). Effect of Zn, Mn, and Fe on Cd accumulation in phytoplankton:
462 Implications for oceanic Cd cycling. *Limnology and Oceanography*, 45(7), 1501–1516.

463 Twining, B. S., & Baines, S. B. (2013). The Trace Metal Composition of Marine Phytoplankton. *Annual*
464 *Review of Marine Science*, 5(1), 191–215. <https://doi.org/10.1146/annurev-marine-121211-172322>

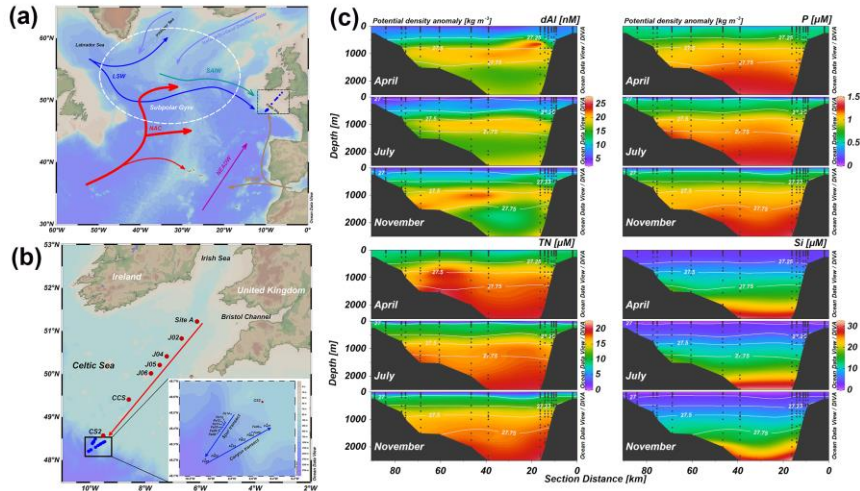
465 Woodward, E. M. S., & Rees, A. P. (2001). Nutrient distributions in an anticyclonic eddy in the northeast
466 Atlantic Ocean, with reference to nanomolar ammonium concentrations. *Deep Sea Research Part II:*
467 *Topical Studies in Oceanography*, 48(4), 775–793. [https://doi.org/10.1016/S0967-0645\(00\)00097-7](https://doi.org/10.1016/S0967-0645(00)00097-7)

468 Wu, J., & Roshan, S. (2015). Cadmium in the North Atlantic: Implication for global cadmium–phosphorus
469 relationship. *Deep Sea Research Part II: Topical Studies in Oceanography*, 116, 226–239.
470 <https://doi.org/10.1016/j.dsr2.2014.11.007>

471 Wyatt, N. J., Milne, A., Woodward, E. M. S., Rees, A. P., Browning, T. J., Bouman, H. A., et al. (2014).
472 Biogeochemical cycling of dissolved zinc along the GEOTRACES South Atlantic transect GA10 at
473 40°S. *Global Biogeochemical Cycles*, 28(1), 44–56. <https://doi.org/10.1002/2013GB004637>

474 Xie, R. C., Galer, S. J., Abouchami, W., Rijkenberg, M. J., De Jong, J., De Baar, H. J., & Andreae, M. O.
475 (2015). The cadmium–phosphate relationship in the western South Atlantic—The importance of mode
476 and intermediate waters on the global systematics. *Marine Chemistry*, 177, 110–123.
477

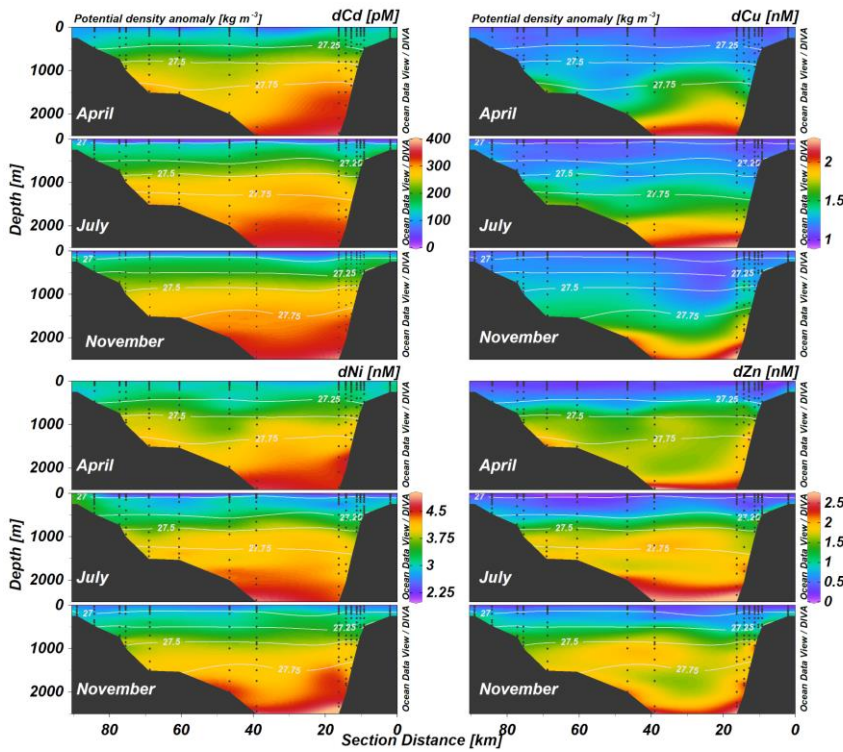
478 **Figures**



479
 480 Fig. 1 (a) The schematic circulation of water masses (NAC: North Atlantic Current; LSW: Labrador Sea
 481 waters; SAIW: Sub-Arctic Intermediate Waters; MOW: Mediterranean Outflow Waters; NEADW:
 482 Northeast Atlantic Deep Waters) in the North Atlantic Ocean; (b) Sampling transects and locations on the
 483 Northeast Atlantic continental margin (Celtic Sea). The red and blue arrows define the shelf and slope
 484 sections, respectively, for Fig. 2 and Fig. 3. (c) Section plots of dissolved aluminum (dAl), phosphate (P),
 485 nitrate+nitrite (TN), and silicic acid (Si) along the slope transect during expeditions in November 2014
 486 (DY018), April 2015 (DY029), and July 2015 (DY033) in Celtic Sea.
 487

Commented [ML68]: Make the black dots larger so can see the sampling resolution

Commented [CX69R68]: Done

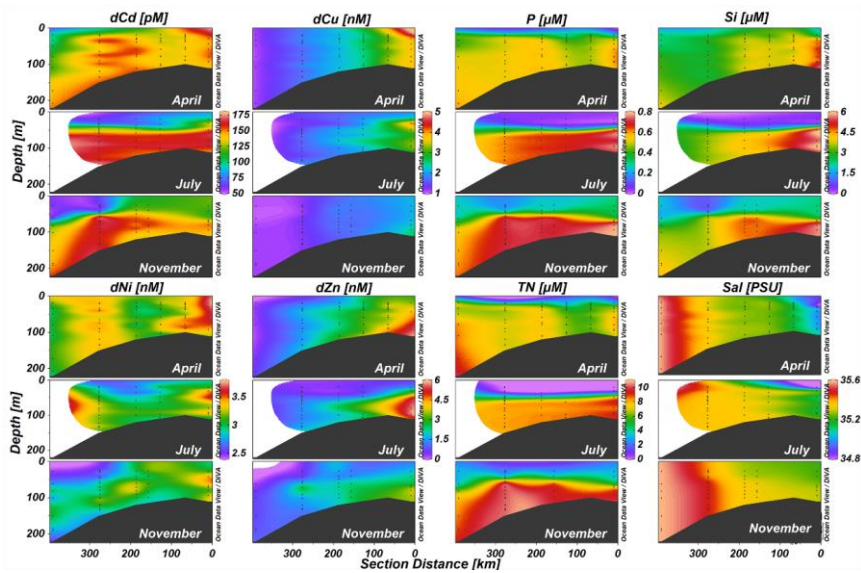


488
 489 **Fig. 2** Section plots of dissolved cadmium (dCd), copper (dCu), nickel (dNi), and zinc (dZn) on the slope
 490 of the Northeast Atlantic continental margin. Samples were taken in November 2014, April 2015, and July
 491 2015, respectively. The section is defined in **Fig. 1b**.
 492

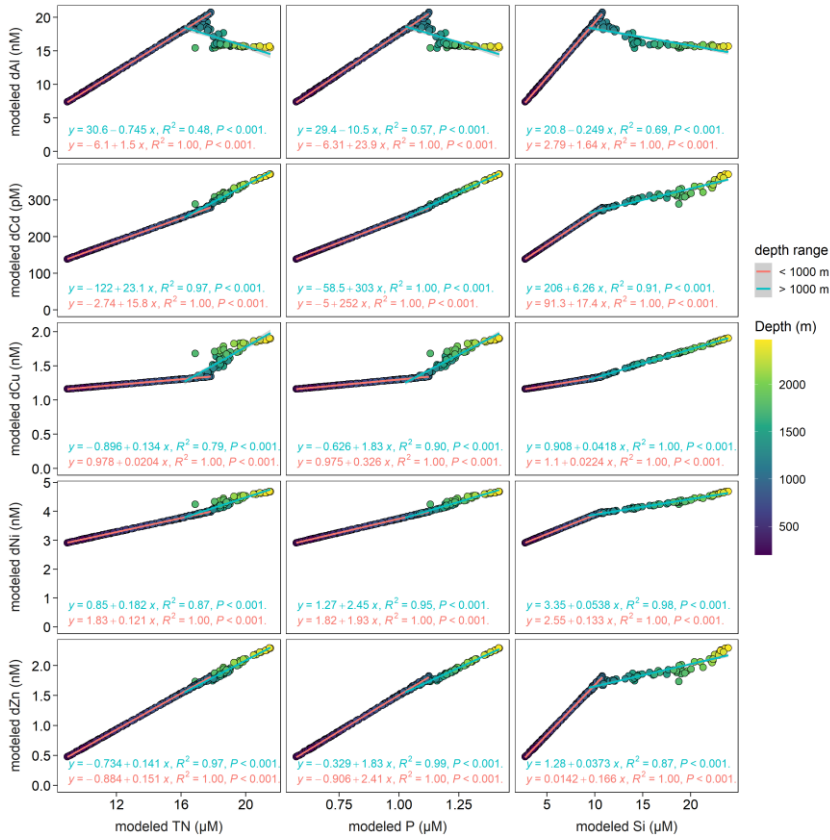
Commented [CX70]: Antony: Some of the surface density isopycnals are a bit squashed, you might want to remove some of the surface ones

It might be better to plot each transect separately? I know this will double the plots but I think at the moment the apparent features in interpolated section in the middle catches the eye, especially in the dCu and dCd. Either that or reduce the interpolation in ODV so that the middle is left blank?

Commented [CX71R70]: The density isopycnals have been modified. I would like to keep the plots with combined transects because we focus on the water mass mixing here.



493
 494 Fig. 3: Section plots of dissolved trace metals (dCd, dCu, dNi, dZn), salinity, and nutrients (nitrate+nitrite
 495 (TN), phosphate (P), silicic acid (Si)) on the continental shelf of the Northeast Atlantic Ocean. The section
 496 is defined in Fig. 1b. Samples were taken in November 2014, April 2015, and July 2015, respectively.
 497



498
 499 **Fig. 4:** Correlations between reconstructed dissolved trace metal (dTM: dAl, dCd, dCu, dNi, dZn) and
 500 nutrient (nitrate+nitrite (TN), phosphate (P), and silicic acid (Si)) concentrations on the Northeast Atlantic
 501 continental slope. Linear regression models were applied to depths < 1000 m (potential density < 27.62 kg
 502 m⁻³) and depths > 1000 m (potential density > 27.62 kg m⁻³), respectively.
 503

Commented [GM72]: You need to remove the grey background

Commented [XC73R72]: done